

Il ciclo profondo del carbonio: rilascio e trasporto di CO₂ dal mantello all'esosfera



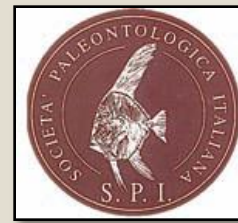
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Dipartimento di Scienze della Terra, Siena



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The long-term carbon cycling



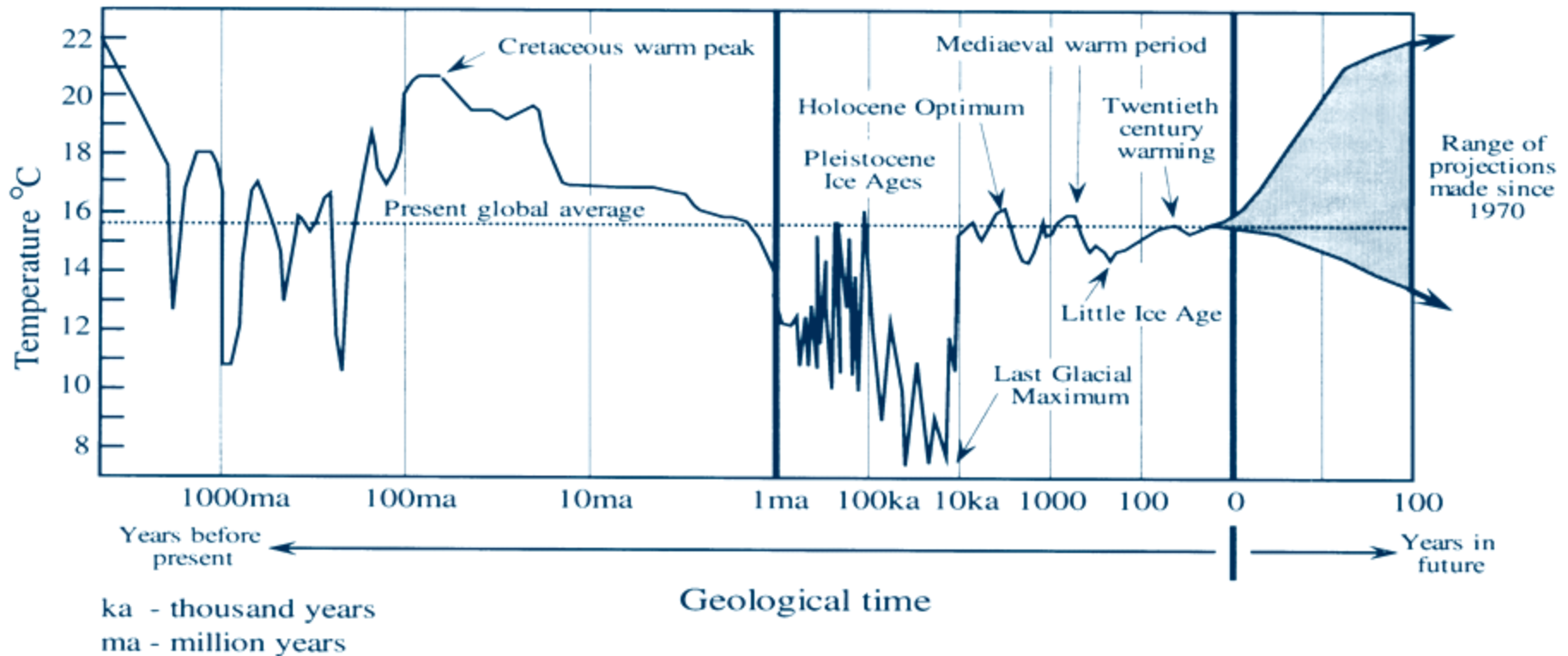
- Introduction – The Earth CO₂ thermostat
- Earth carbon emissions and reservoirs
- Cycling of carbon in subduction zones
- Modelling deep carbon cycling

Long term Carbon cycling

The Earth thermostat



- CO₂ 260 ppm 10,000 yr ago
- CO₂ 280 ppm in 1750



Sagan, & Mullen, 1972 Science ; Walker, et al, 1982, *JGR.*; Zeebe, & Caldeira, 2008 Nature G;
Lovelock, 1974, *Tellus*; Bryant, 1997; Archer, 2008, Nature G.

The perturbation

Anthropocene CO₂ increase

Carbon Dioxide

Anthropogenic emissions

Earth emissions

CO₂ level increase from 1750-2005

Conversion Rate

T increase of in the last 250 years

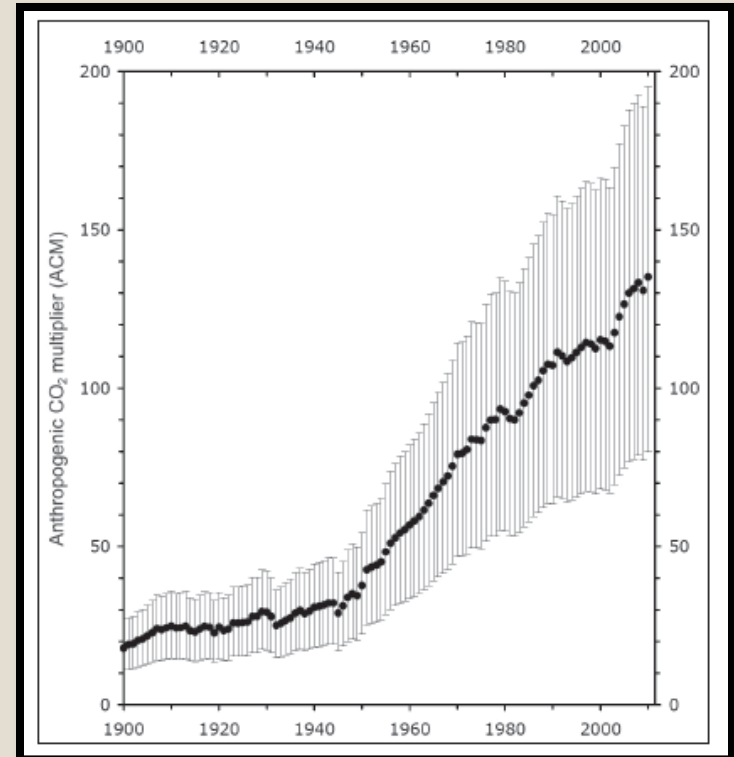
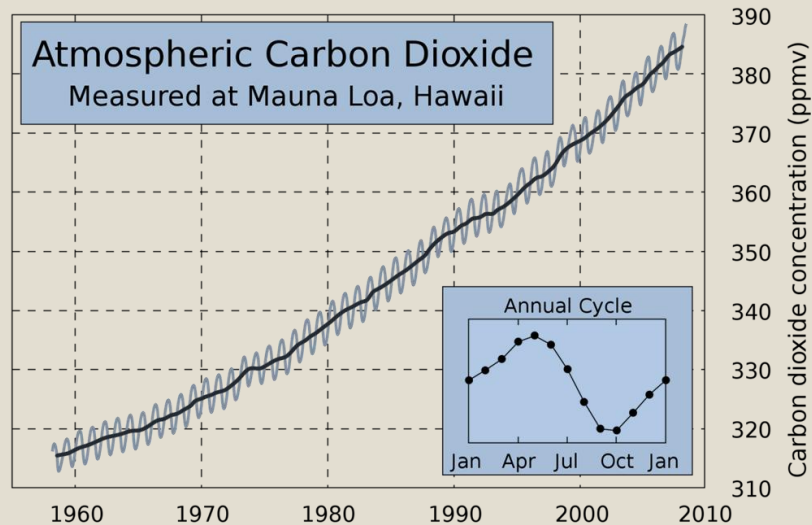
35 GtCO₂/yr or 10 Gt C/y

0.26 GtCO₂/yr (0.18-0.44Gt) or 0.1 Gt C/y

380-280=100 ppm

1 ppm= 2.12 GtC and 7.78 GtCO₂

0.5 °C



Long-term (stable) carbon cycle



- “The irrelevance of the long-term carbon cycle to the immediate future is an important point, since the global warming climate event will last for as long as it takes these slow processes to act”

- ...”The geologic carbon cycle sets the stage for for carbon cycle gymnastic on faster and perhaps more immediately relevant time scales”...

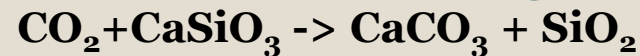
- ...”**On geologic time scales the carbon cycle acts to stabilize the climate of the Earth in mysterious and wondrous ways**”...

The long-term Earth carbon cycle

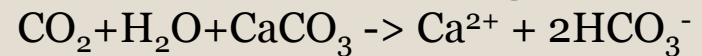
Main Reactions

- **$\text{CaSiO}_3 + \text{CO}_2 \leftrightarrow \text{CaCO}_3 + \text{SiO}_2$**
 - (UREY, 1952)

Silicate rock weathering



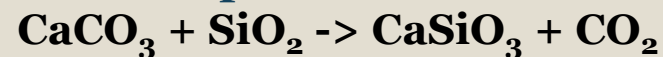
Carbonate rock weathering



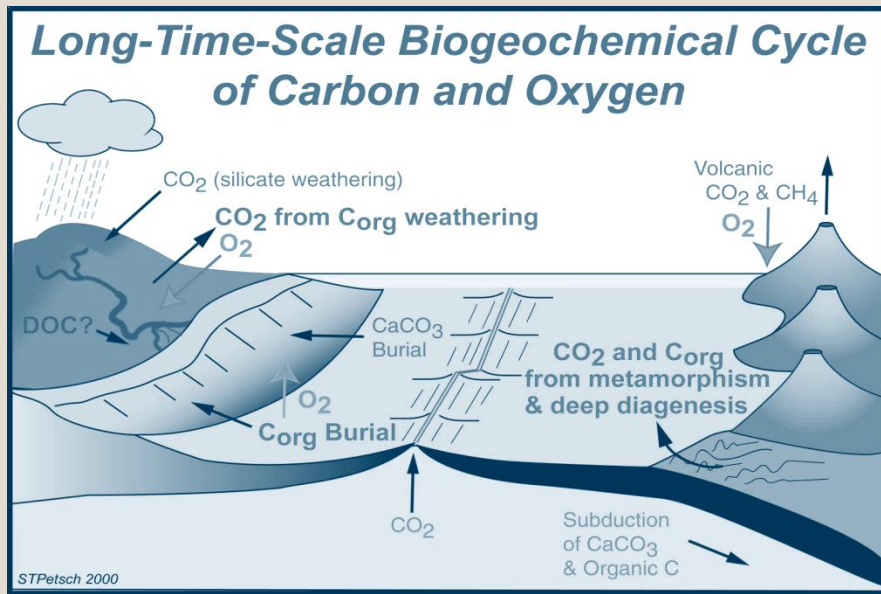
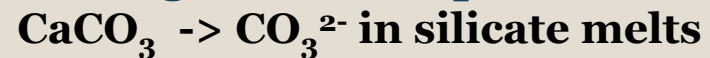
Carbonate formation in oceans



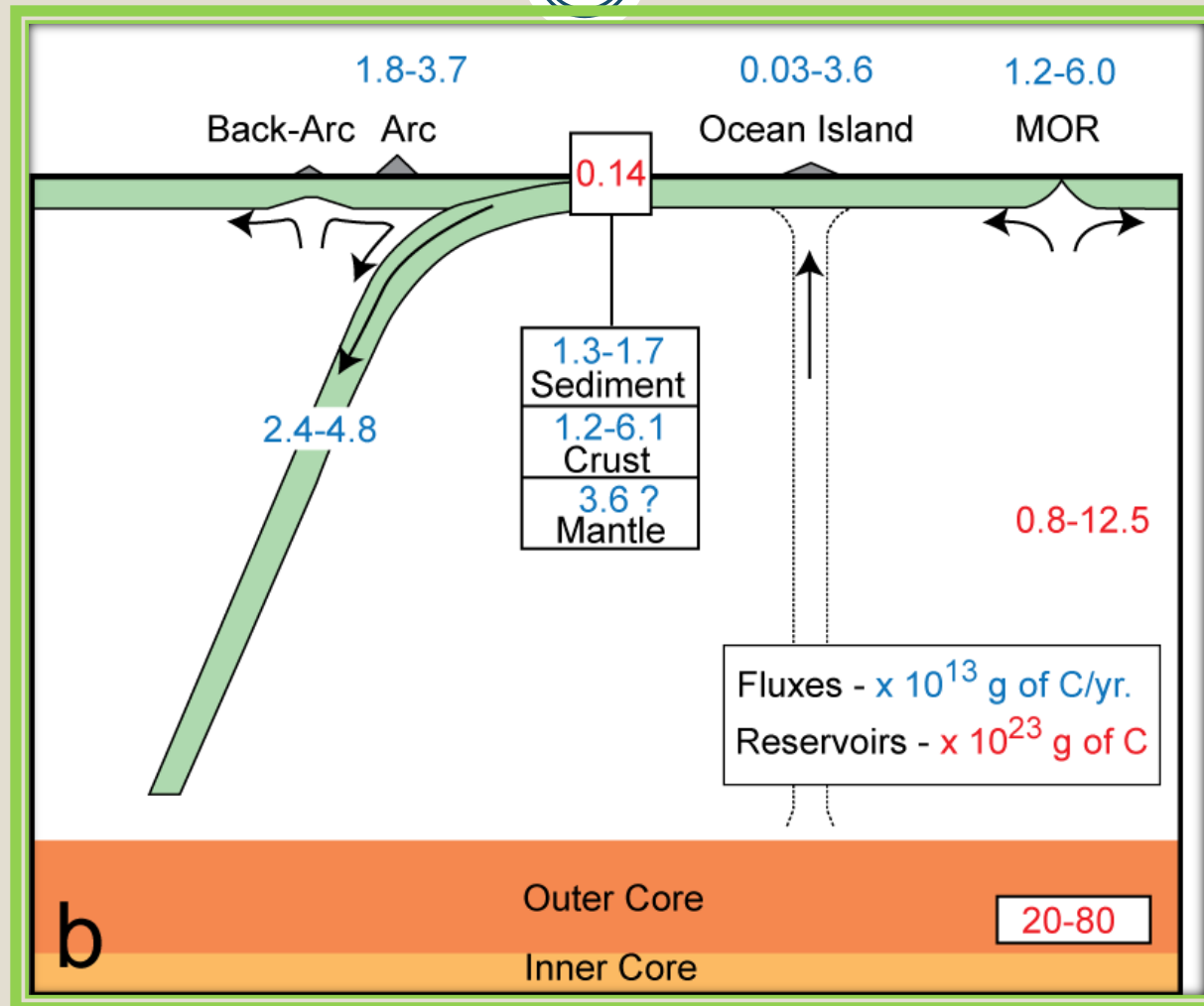
Metamorphic breakdown of carbonates



Melting of metamorphic carbonates



Earth Carbon reservoirs and fluxes



Earth CO₂ degassing

Point source measurements



Earth CO₂ degassing

1970  1992

VOLCANIC CO ₂	SUB AERIAL	SUB AQUEOUS	Total
	Mt/y	Mt/y	
Regional Measurements			
Etna crater 1976 - 1985 (Italy) <i>Allard, 1991</i>	25.5		
Augustine crater (Alaska, USA) <i>Gerlach 1991</i>	2.2		
Mt. St. Helens crater (Washington, USA) <i>Gerlach 1991</i>	1.8		
Nyiragongo crater (Zaire) <i>Le Guern, 1987</i>	1.2		
White Island crater (New Zealand) <i>Rose et al., 1986</i>	0.66		
Kilauea crater (Hawaii, USA) <i>Greenland et al., 1985</i>	0.47		
Vulcano crater (Italy) <i>Baubron et al., 1990</i>	0.06		
Erta 'Ale (Ethiopia) <i>Le Guern et al., 1979</i>	0.02		
Mt. St. Helens 1980 eruption (USA) <i>Casadevall et al., 1983</i>	4.8 - 22		
Lake Nyos 1986 event (Cameroon) <i>Kousakabe et al., 1989</i>	1.24		
Measurements - Total 1970 - 1992			31.9
Measurements - Average 1992 (n=8)			4
Global Extrapolations*			
<i>Le Guern, 1982</i>	48		
<i>Gerlach, 1991</i>	80	22 - 40	102 - 120
<i>Allard, 1992</i>	66		
<i>Marty and Le Cloarec, 1992</i>	66 - 110		
Global Estimates[§]			
<i>Williams et al., 1992</i>	64		
<i>Varekamp et al., 1992</i>	145	66 - 97	
Range Global, 1970 - 1992	48 - 145	22 - 97	70 - 242

VOLCANIC CO ₂	SUB AERIAL	SUB AQUEOUS
	Mt/y	Mt/y
Regional Measurements		
Pinatubo 1991 eruption (Philippines) <i>Gerlach et al., 1996</i>	42	
Popocatepetl crater (Mexico) <i>Varley and Armienta, 2001</i>	21.9	
Etna crater 1993 - 1997 (Italy) <i>Allard, 1998</i>	4 - 13	
Masaya Caldera diffuse (Nicaragua) <i>Pérez et al., 2000</i>	10.5	
Etna flanks (Italy) <i>D'Alessandro et al., 1998</i>	0.6 - 6	
Oldoinyo Lengai crater (Tanzania) <i>Koepnick, 1995</i>	2.2 - 2.6	
Stromboli crater (Italy) <i>Allard et al., 1994</i>	1 - 2	
Masaya crater (Nicaragua) <i>Burton et al., 2000</i>	0.8 - 1.13	
Cerro Negro total (Nicaragua) <i>Salazar et al., 2001</i>	1	
Rabaul Caldera (Papua N.G.) <i>Perez et al., 1998</i>	0.88	
Erebus crater (Antarctica) <i>Wardell and Kyle, 2003</i>	0.7	
Redoubt crater (Alaska, USA), <i>Casadevall et al., 1994</i>	0.7	
Canadas Caldera (Tenerife, Spain) <i>Hernandez et al., 2000a</i>	0.2	
Teide flanks (Tenerife, Spain) <i>Salazar et al., 1997</i>	0.2	
Vulcano fumaroles (Italy) <i>Italiano et al., 1998</i>	0.13	
Usu crater (Japan) <i>Hernandez et al., 2001</i>	0.04 - 0.12	
Mout Baker crater (Washington, USA) <i>Mc Gee et al., 2001</i>	0.07	
Stromboli flanks (Italy) <i>Carapezza and Federico, 2000</i>	0.07 - 0.09	
Hakkoda flanks (Japan) <i>Hernandez et al., 2000b</i>	0.04	
Measurements - Total 1992 - 2002	83 - 90.3	
Measurements - Total 1970 - 2002	114.9 - 122.2	
Average 1970 - 2002 (n=24)	4.6 - 4.7	

NON-VOLCANIC CO ₂	SOIL EMISSION
	Mt/y (km ²)
Regional Measurements	
Back arc Pacific Rim, <i>Seward and Kerrick, 1996</i>	44
Central Apennine (Italy) <i>Chiodini et al., 2000</i>	4 - 13 (12,000)
Larderello and Amiata (Italy) <i>Chiodini et al. 2000</i>	2.2
Indonesia - Philippines, <i>Seward and Kerrick, 1996</i>	1.8
Nisyros soil (Greece) <i>Brombach et al., 2001</i>	0.8
Siena basin (Italy) <i>Etiopo et al., 1996</i>	0.5 (200)
Taupo volcanic zone (N Zealand) <i>Seward and Kerrick, 1996</i>	0.44
Cascade Range (Oregon, USA) <i>James et al., 1999</i>	0.4
St. Andreas fault (California, USA) <i>Lewicki and Brantley, 2000</i>	0.3 (60)
Ustica (Italy) <i>Etiopo et al., 1999</i>	0.26 (9)
Mammoth Mountain (California, USA) <i>Rahn et al., 1996</i>	0.15 (0.4)
Yellowstone (Wyoming, USA) <i>Werner et al., 2000</i>	0.18 (3.5)
Measurements - Total 1992 - 2002	55 - 64

1992 - 2010

Global Extrapolations*			Total
Brantley and Koepnick, 1995	88 - 132		
Kerrick, 2001	88 - 110		
Morner and Etiopo, 2002		300	> 600
Kerrick and Caldeira, 1998			44
Global Estimates[§]			
Sano and Williams, 1996	136	44 - 132	
Marty and Tolstikhin, 1998	242	57 - 136	367
Range, 2002	110 - 242	44 - 136	600 367 - 900

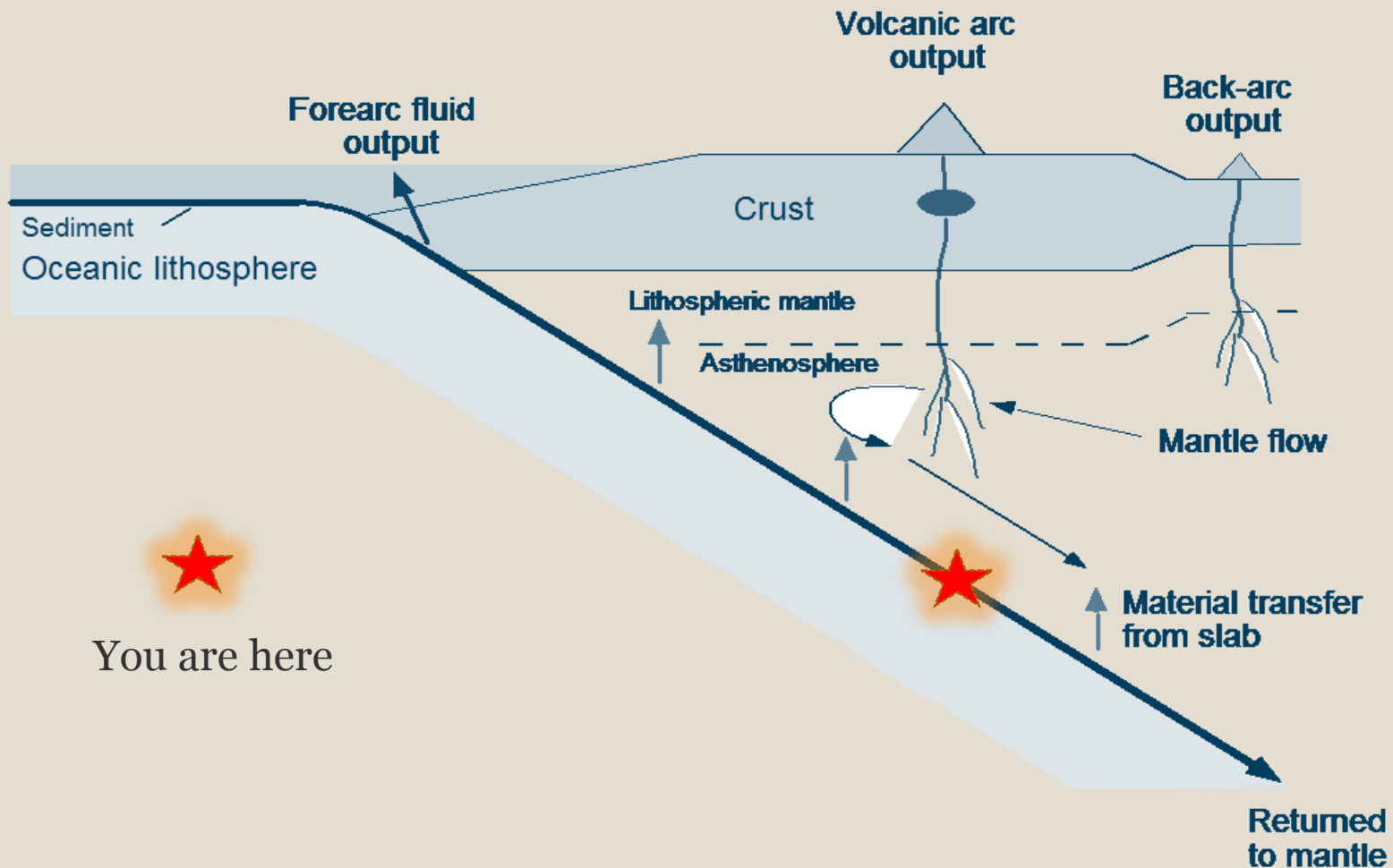
Measurements of natural point source emissions



- From a total atmospheric CO₂ of 0.01 % in the 1990's up to 0.1 % in the 2000's of the anthropogenic emissions.
- the CO₂ emission rates from single volcanoes can be highly variables, differing of one or two orders of magnitude.
- Such a constantly increasing trend in a frame were the degassing processes are unknown indicates that the natural Earth's emissions are unquantified.
- The reasons behind variations between different volcanoes are unknown.

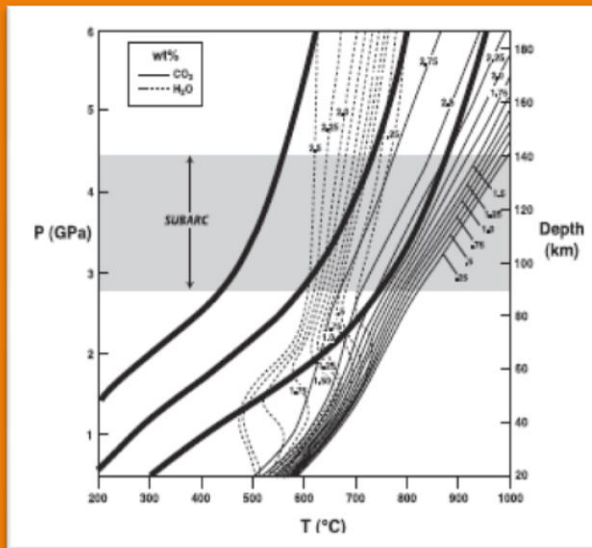
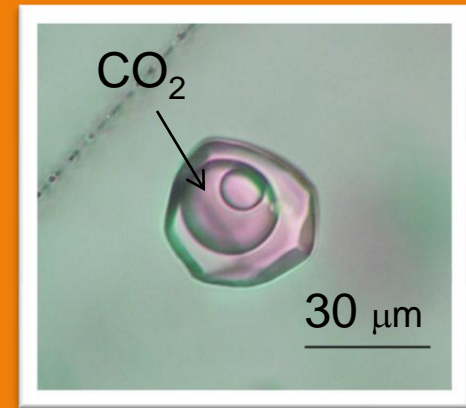
From Volcanology to Petrology

Deep cycling of Carbon



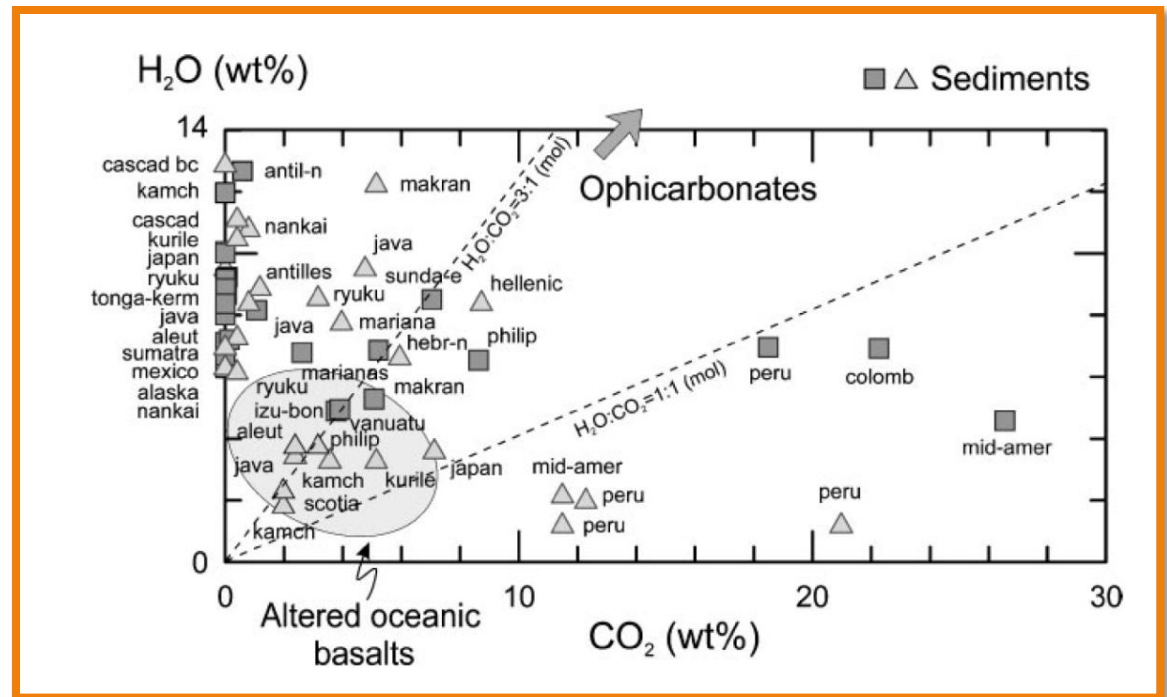
How do we model C fluxes during subduction?

- Fluid inclusions
- HP metamorphic rocks
- Experimental petrology

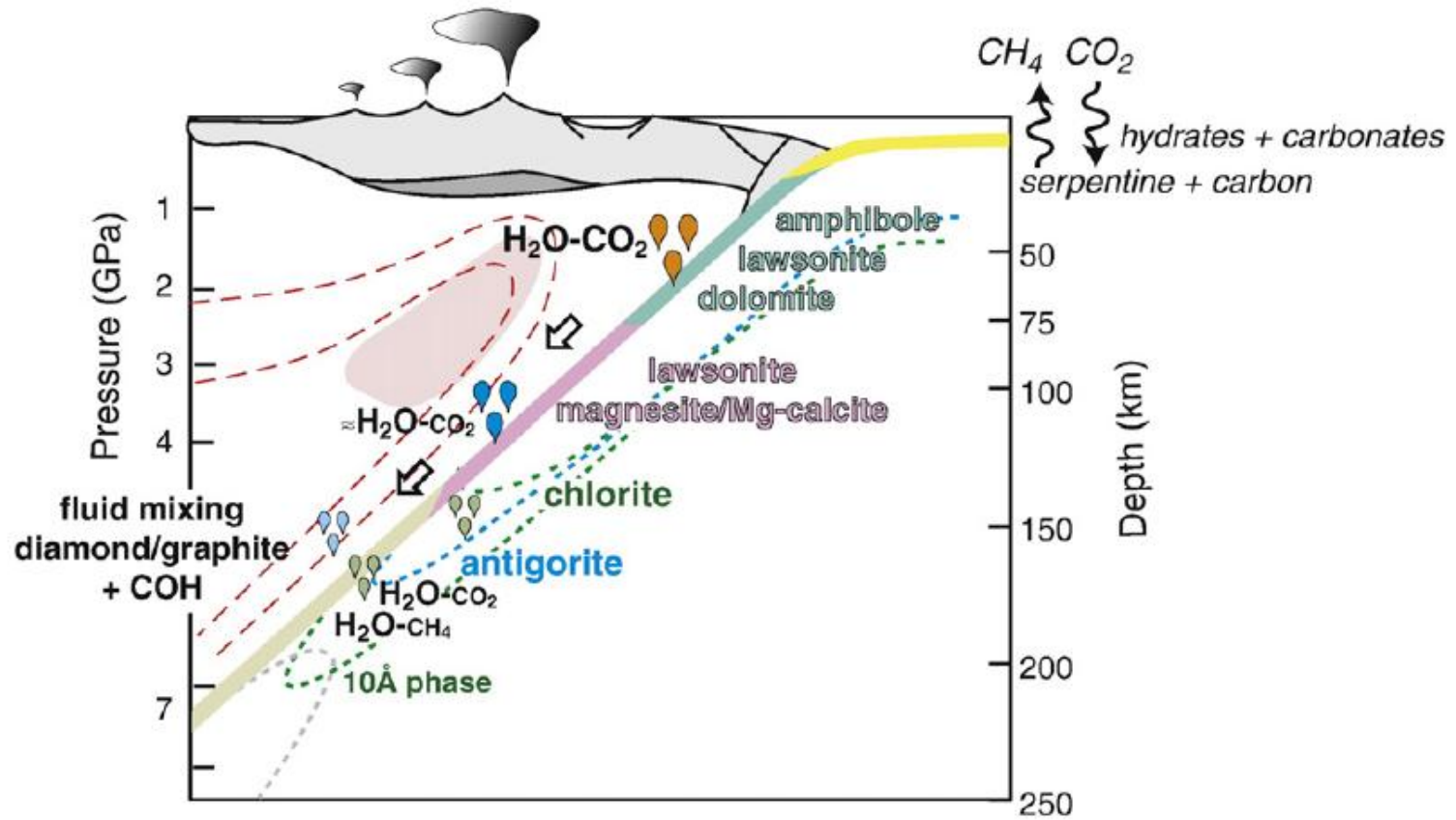


Subducted Carbon

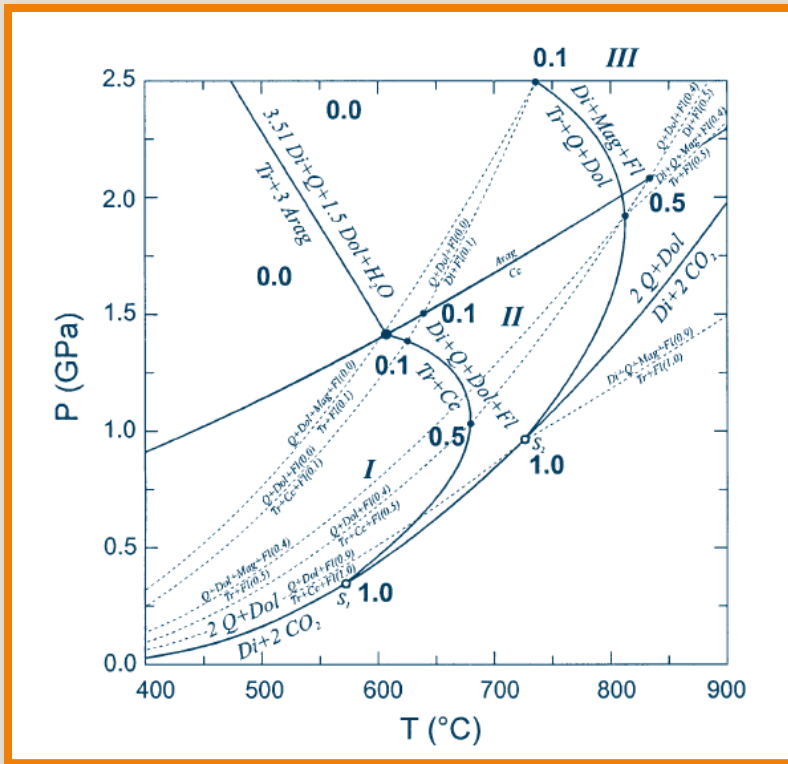
- Oceanic metabasalts (41 x Mt/yr)
- Ocean floor sediments (24 x Mt/yr)



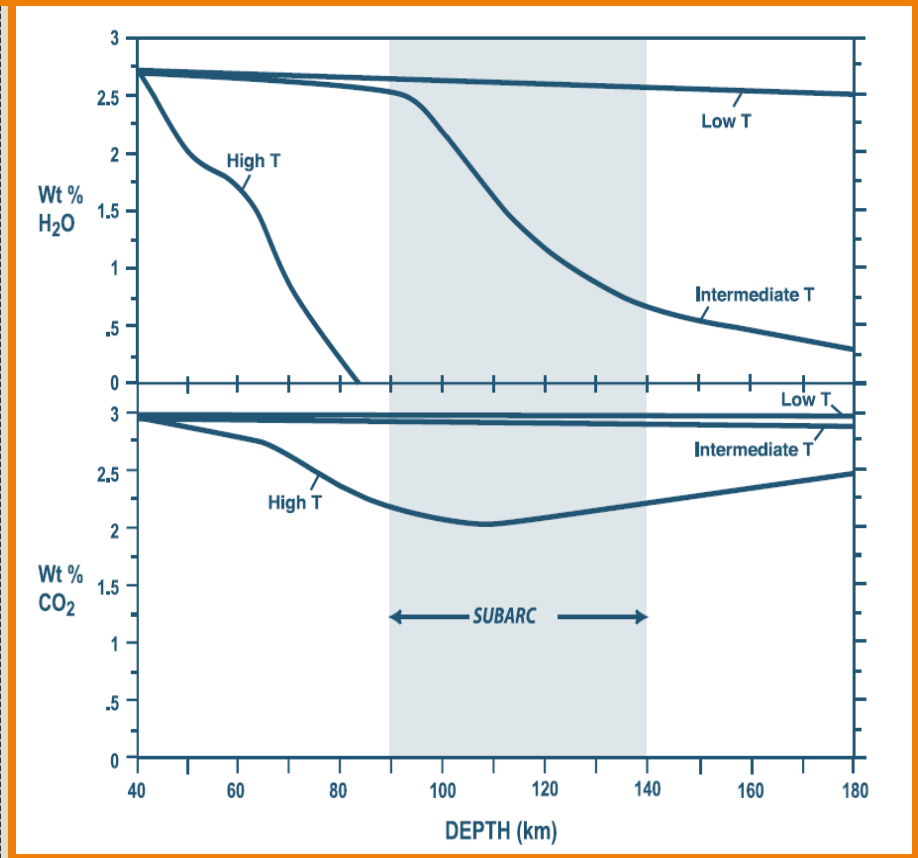
Carbon release from slab



Metamorphic breakdown of carbonates



Molina & Poli, 2000 EPSL



Kerrick & Connolly, 2001 EPSL

Subduction -Volcanic degassing CO₂ imbalance



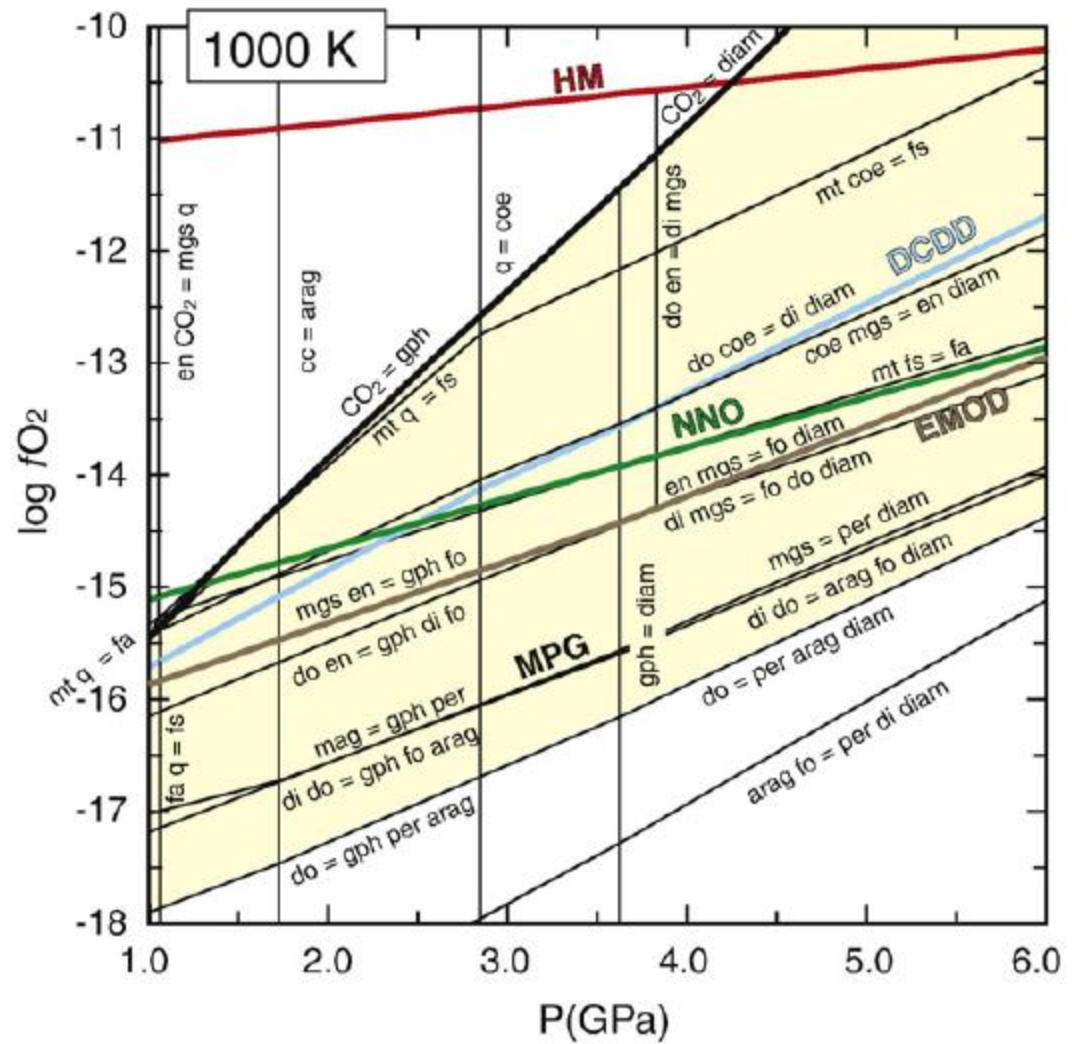
Decarbonation of the subducted slab is generally limited, since it requires higher T than most subduction slabs.

HOWEVER

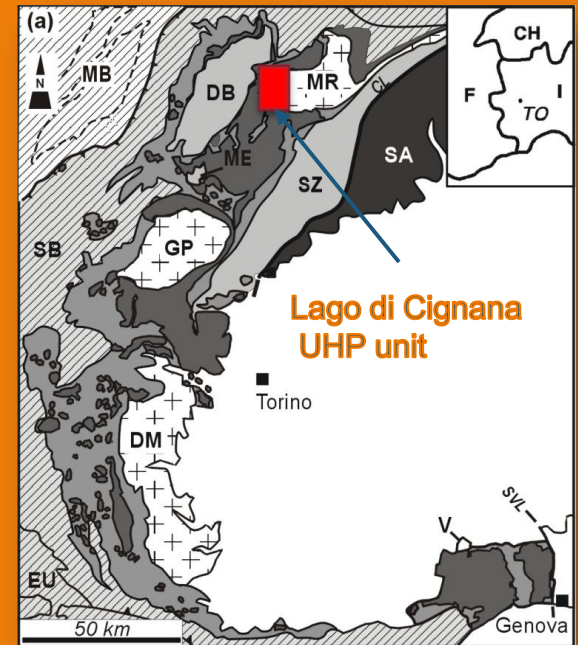
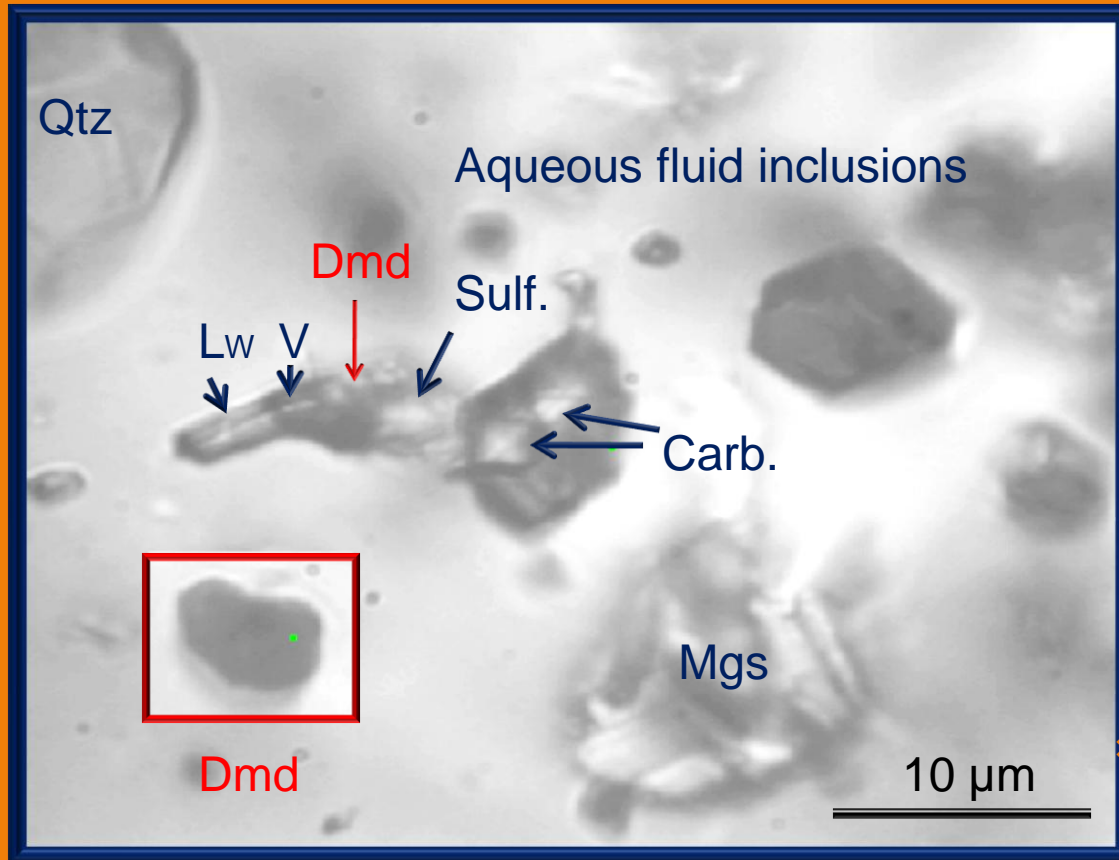
Observations of CO₂ flux and inferences based on CO₂/³He in volcanic gases lead to the conclusion that about half of the CO₂ flux from volcanoes comes from arc volcanism, and that 80% of this is recycled from subducted materials.

Marty and Tolstikhim (1998) CG; Molina and Poli (1998) EPSL; Kerrick and Connolly (2001) N; Sano and Williams EPSL (1996)

fO₂ control on Carbon speciation during subduction



Microdiamonds in UHP rocks (W Alps)



Segment of former oceanic crust

Eclogite facies metabasics and metasediments

Most diamond inclusions $< 10 \mu\text{m}$

Most diamonds within fluid inclusions $\leq 1 \mu\text{m}$

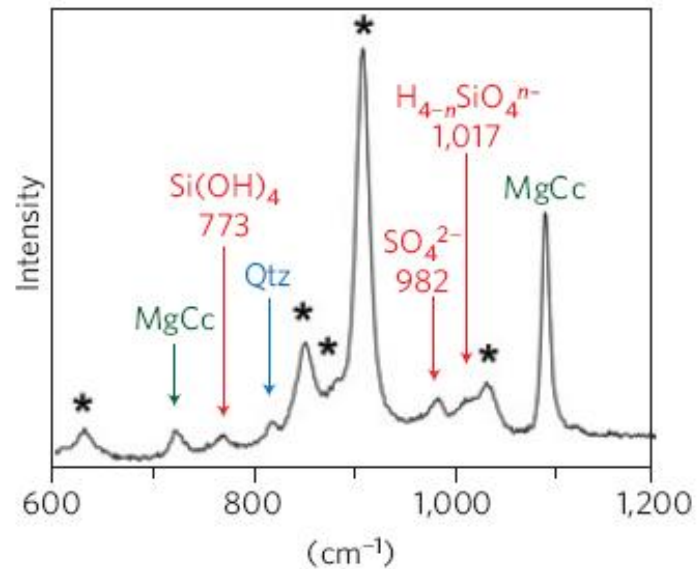
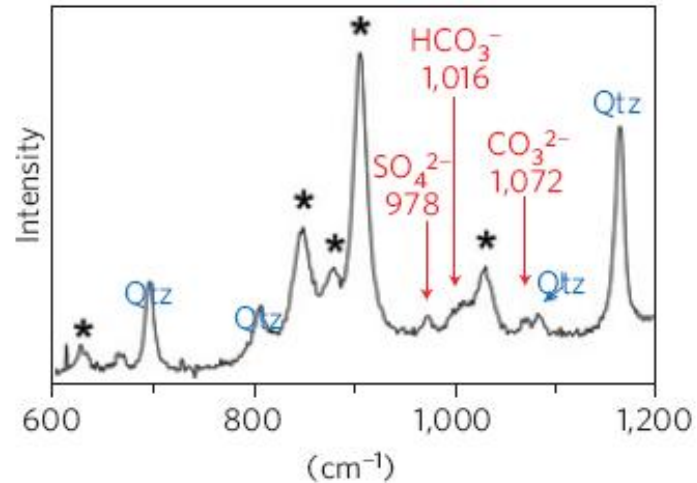
Subduction fluids

$\text{HCO}_3^- > 0.016 \text{ mol/kg H}_2\text{O}$

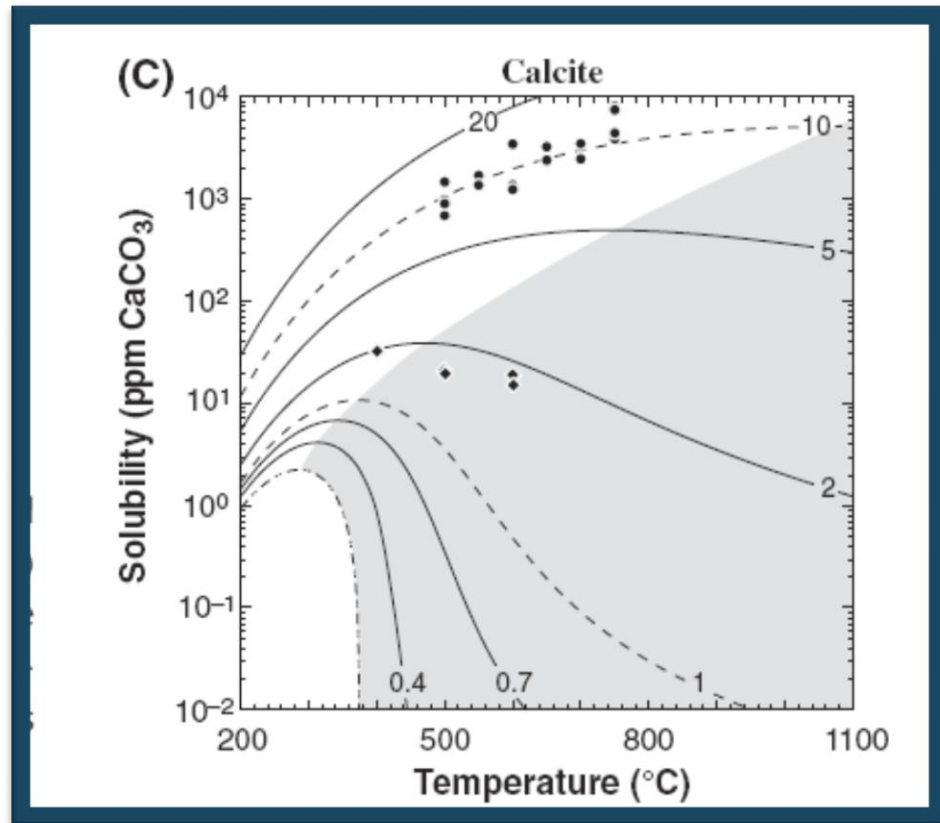
$\text{SO}_4^{2-} > 0.002 \text{ mol/kg H}_2\text{O}$

$\text{CO}_3^{2-} > 0.006 \text{ mol/kg H}_2\text{O}$

Raman spectra collected in liquid water contained within fluid inclusions
Excitation spot $1 \times 1 \times 5 \mu\text{m}$



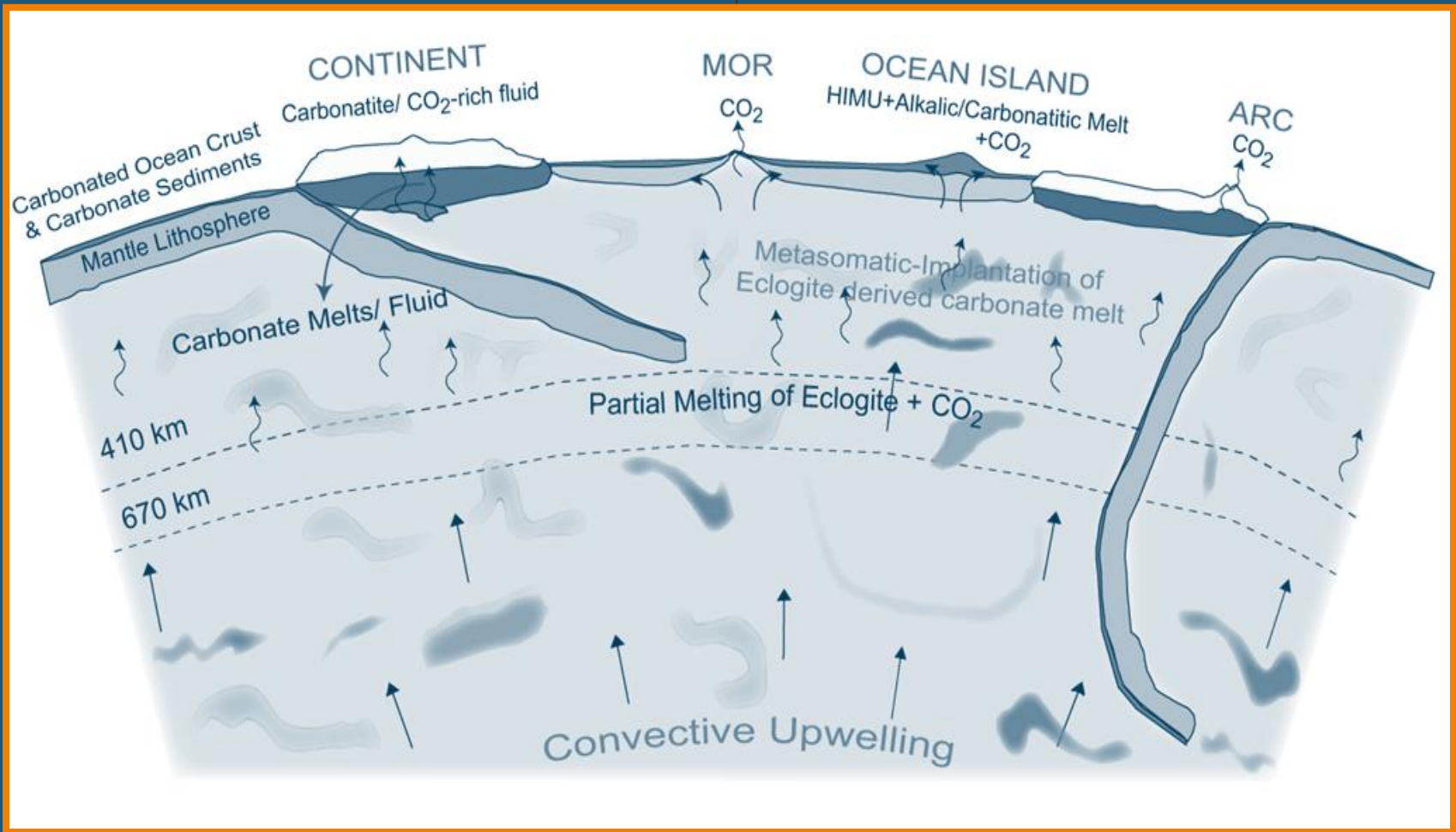
carbon release from the slab through dissolution



Subduction fluids

($T= 600^{\circ}\text{C}$; $P= 3.4 \text{ GPa}$, or 120 km)

Deep cycling of Carbon



Mantle composition



- Scale lengths of heterogeneities (metasomatism)
- Does recycled oceanic crust remain distinct & for how long
- Chemical modifications introduced by plate subduction plus delamination of continental lithospheric roots
- Role of CO₂ v H₂O in mantle melting

Morb - OIB



Present-day mantle reservoirs The MORB-OIB spectrum

Virtually all the geochemical end-members can be explained by an origin from mantle sources that have experienced interaction with crustal “material” introduced at depth.

HIMU

recycled subducted
oceanic crust (MORB)

EMI

recycled subducted
pelagic sediment

EM II

recycled subducted
terrigenous sediment

DM

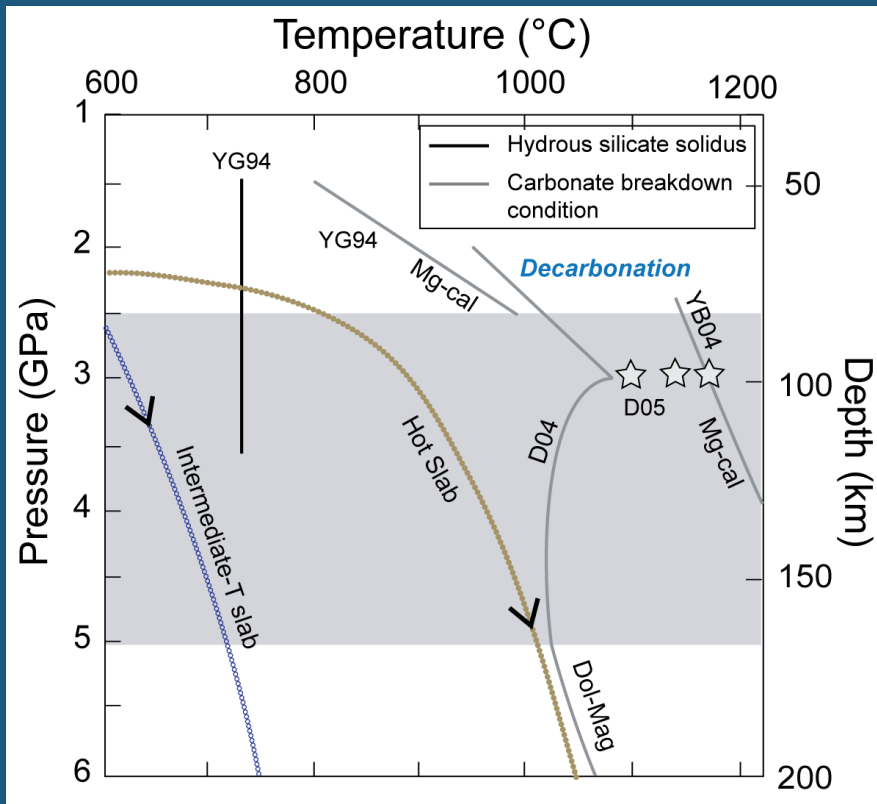
strongly depleted
convecting upper mantle

FOZO

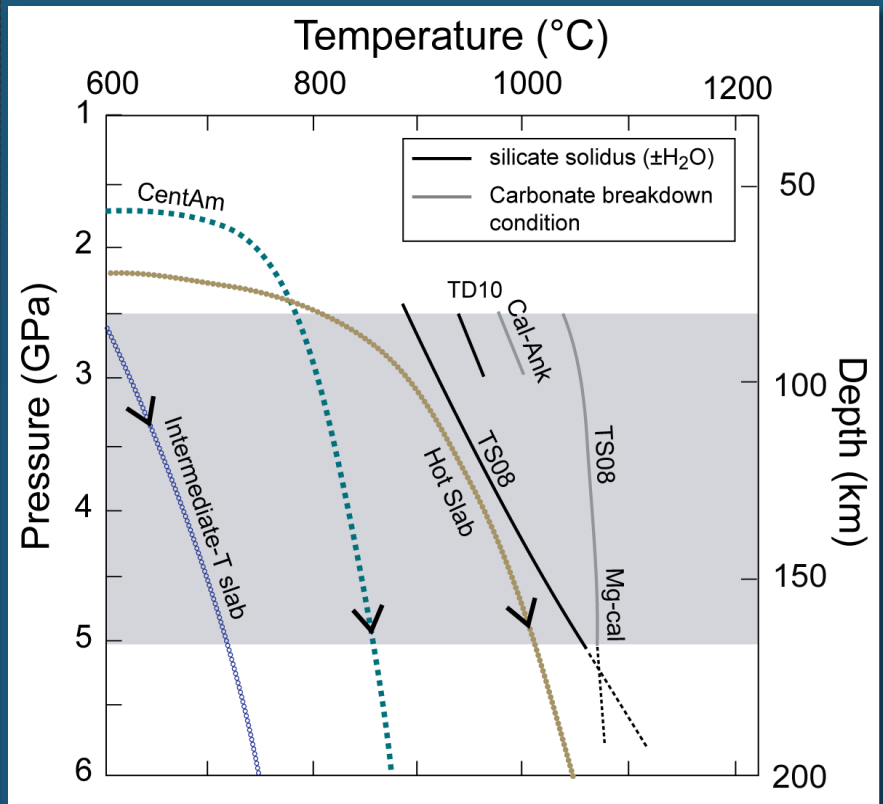
? plume source
? lower mantle

Melting of carbonated rocks

Carbonated basalts

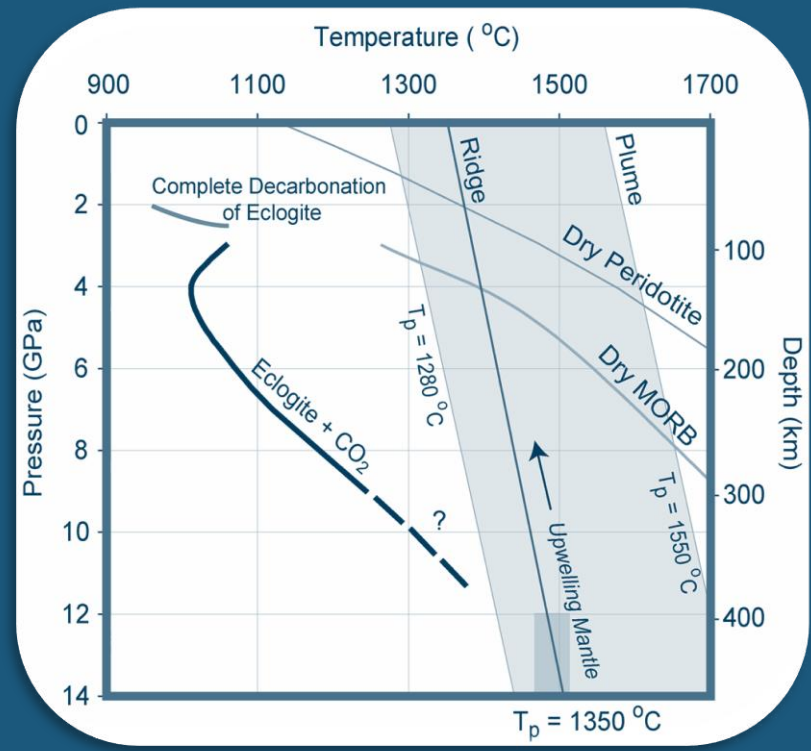
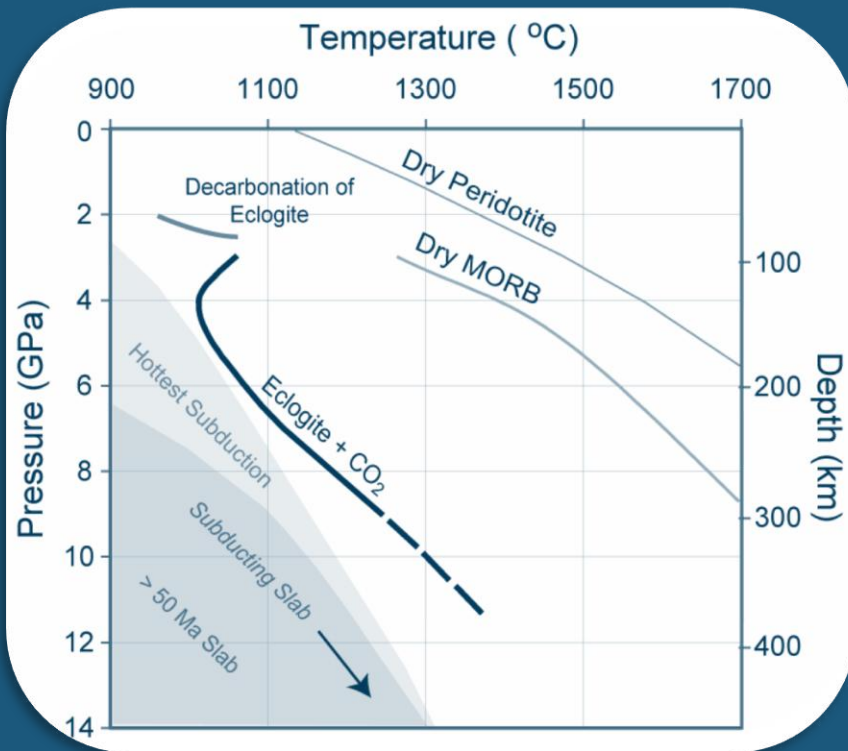


Carbonated sediments

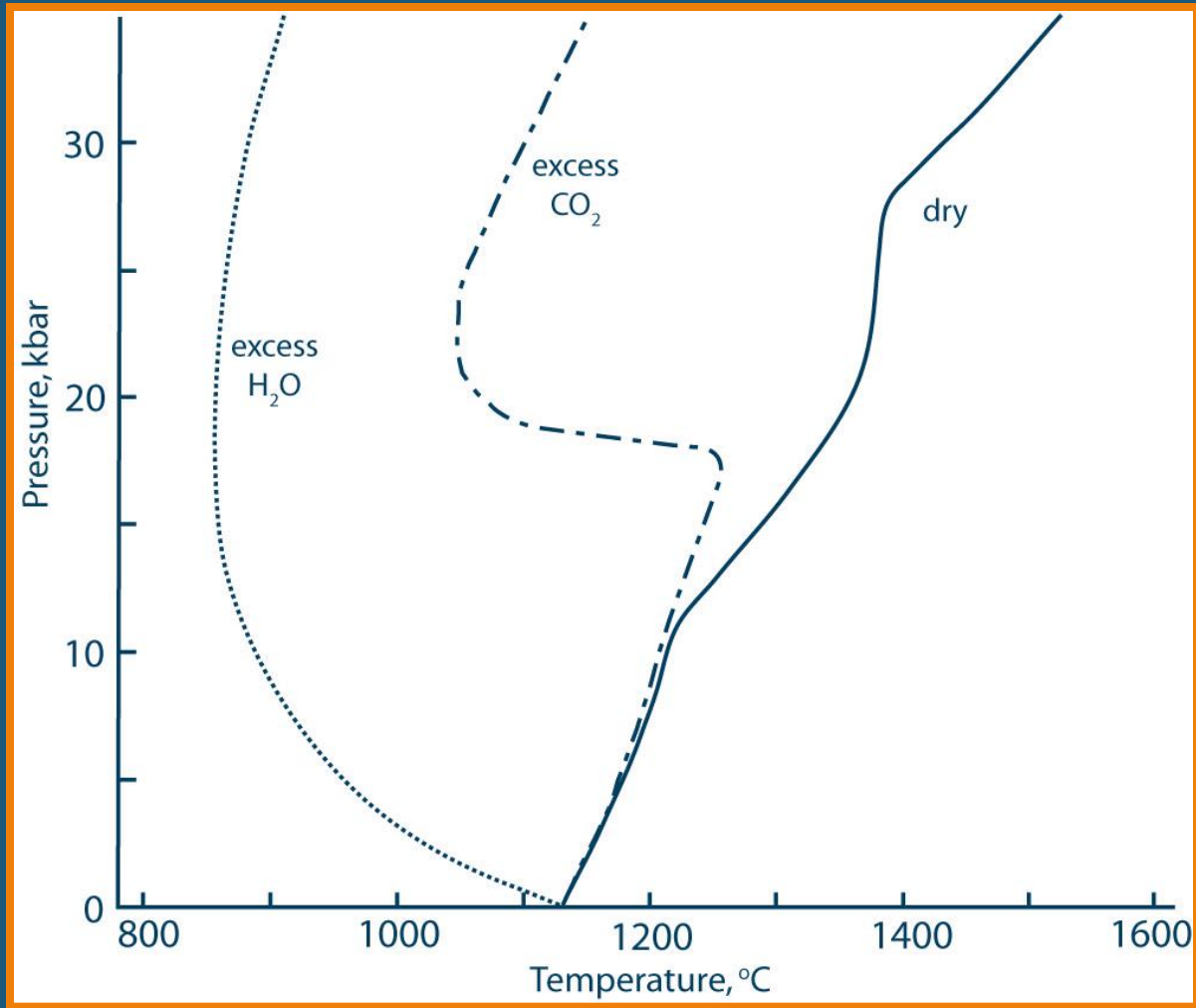


(Yaxley and Green, 1994; Yaxley and Brey, 2004; Tsuno and Dasgupta, 2010; Dasgupta et al., 2004, 2005; Thomsen and Schmidt, 2008;

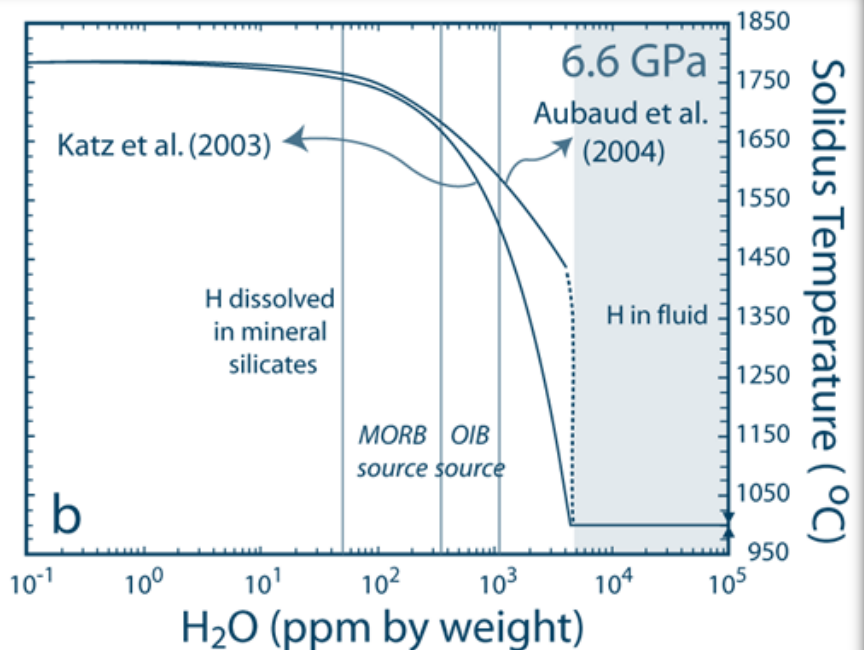
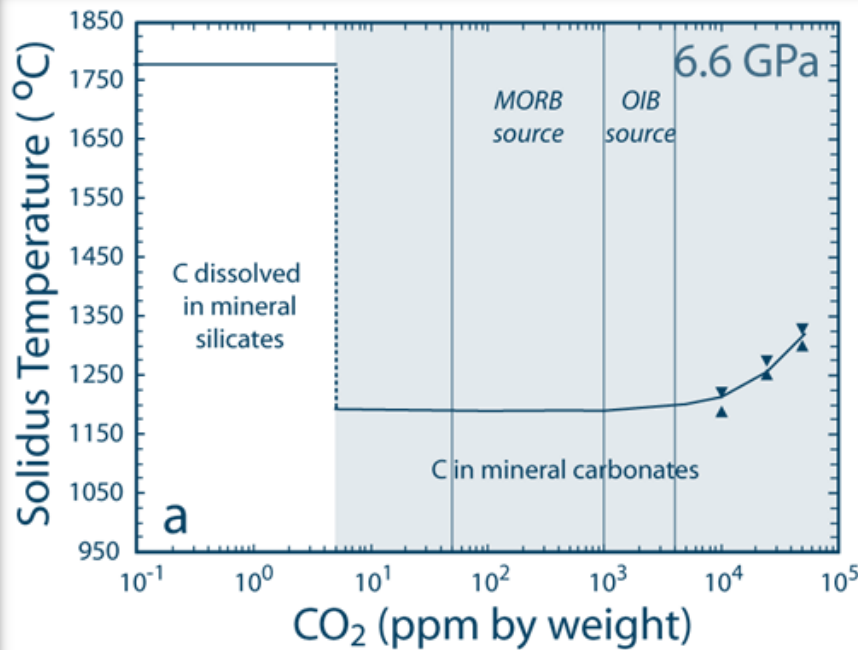
Melting of carbonated eclogite



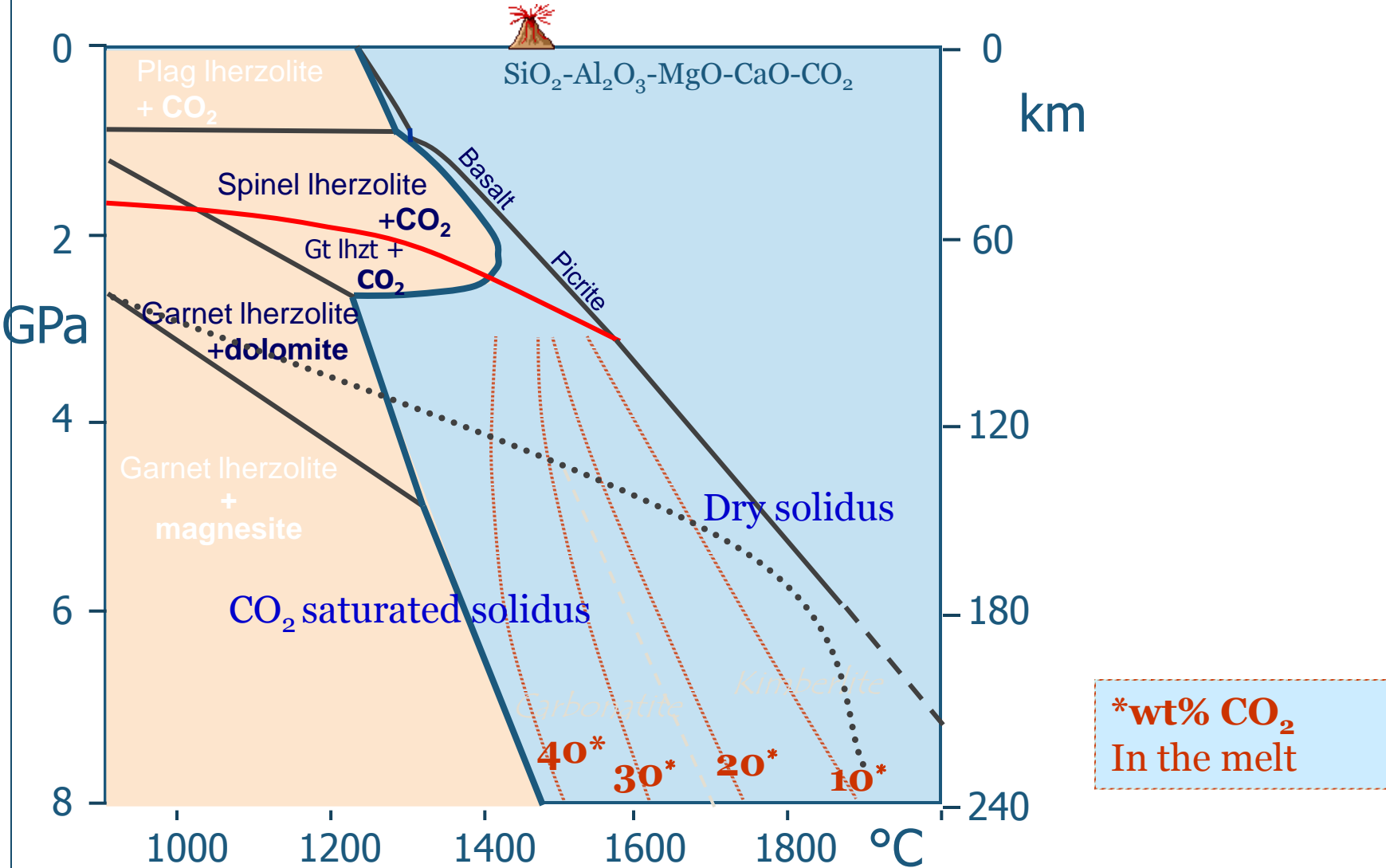
Peridotite solidus



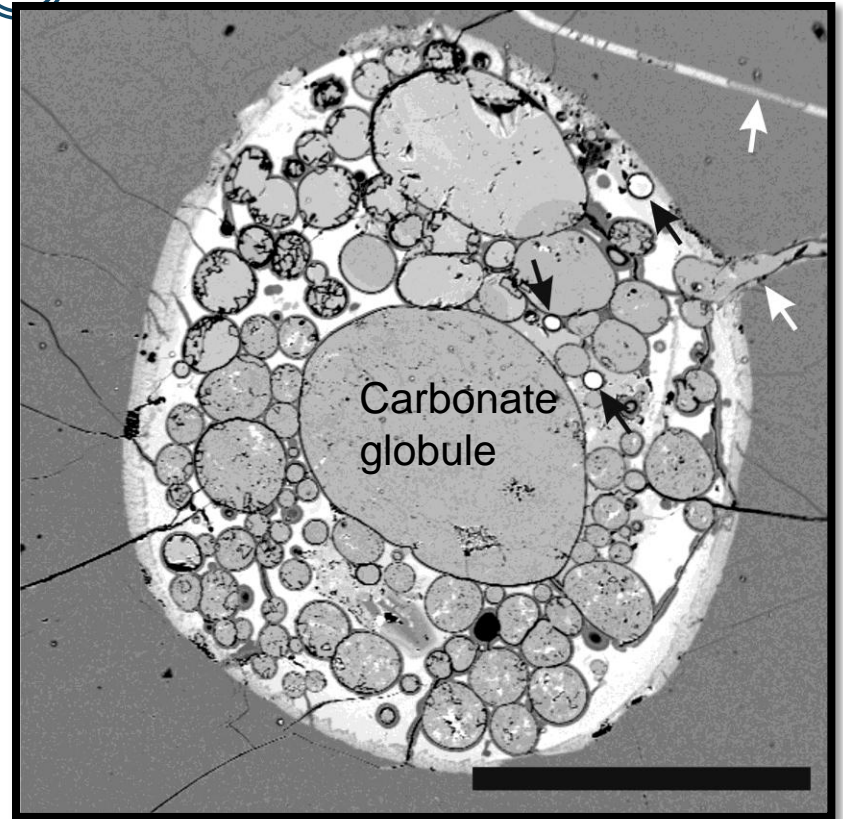
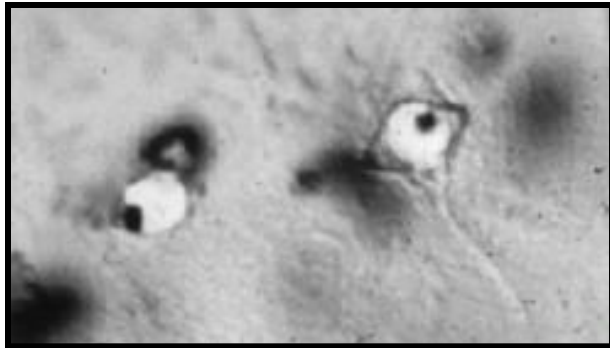
Peridotite Solidus – H₂O and CO₂



Transition from carbonate melts to carbonated silicate melts



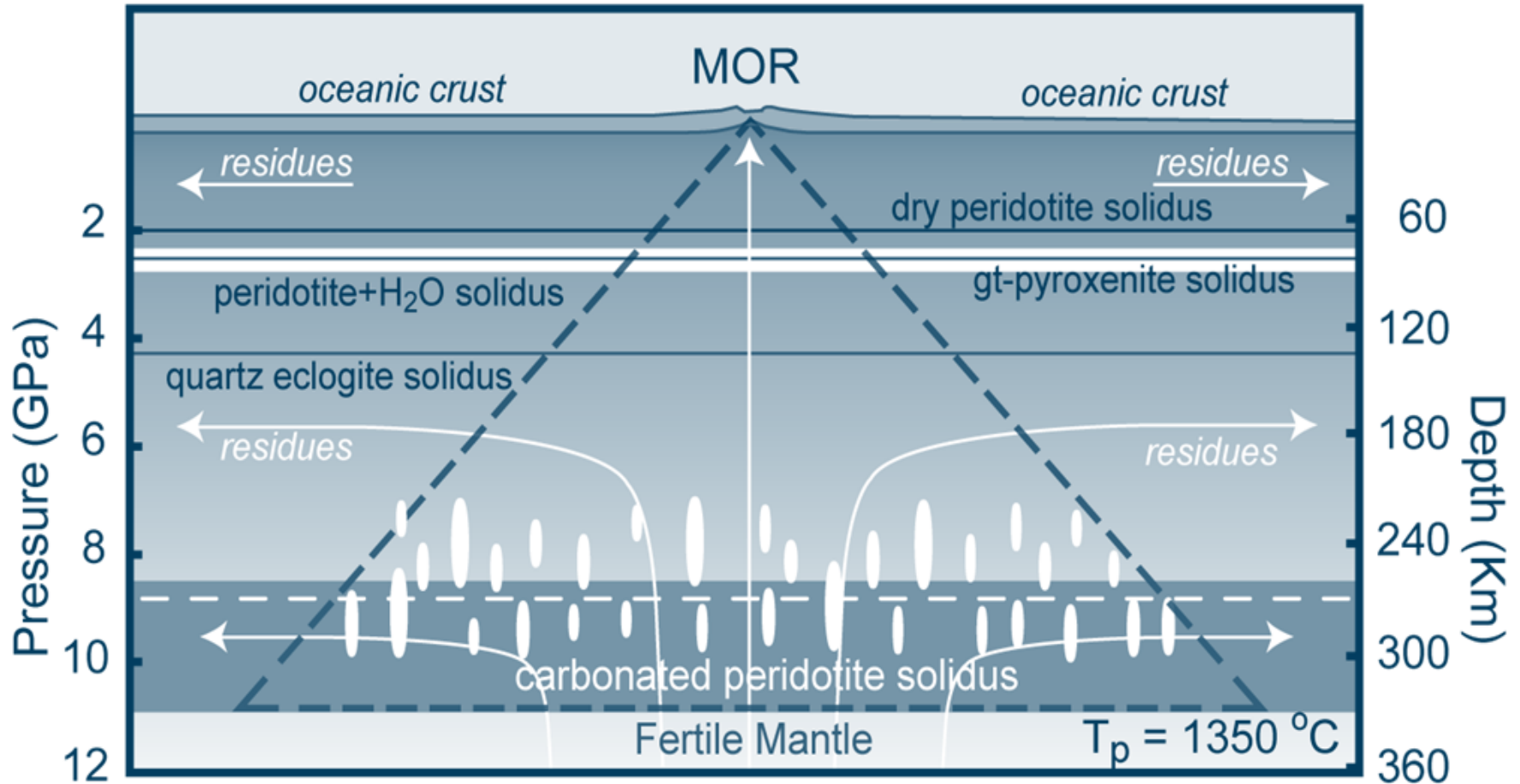
Carbonate globules in mantle rocks



V. Hurai

100 μm

Fluxes of mantle CO₂ between 120 - 3,400 Mt/yr beneath mid ocean ridges (MOR).



Deep Carbon cycling beneath Italy

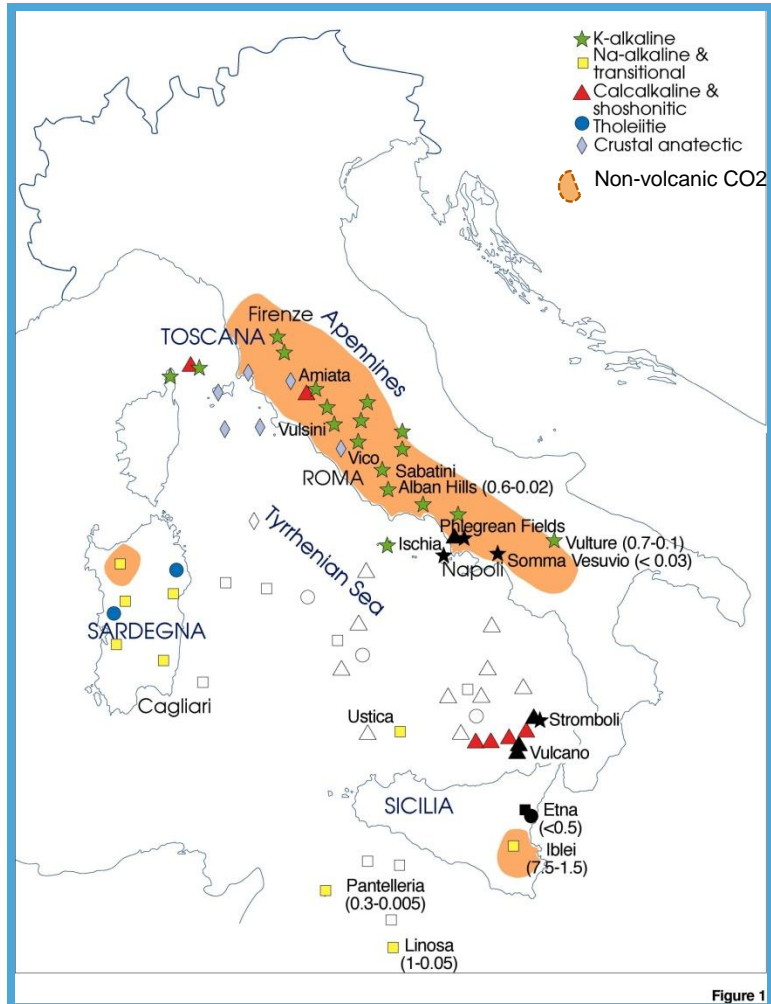


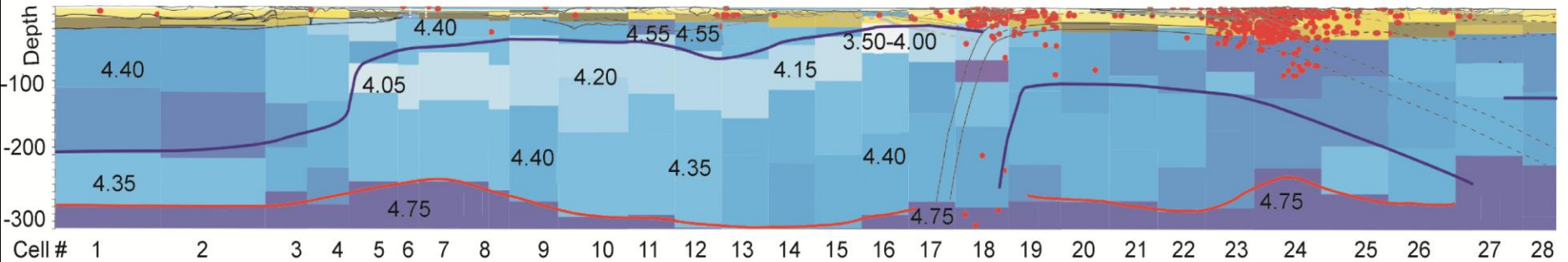
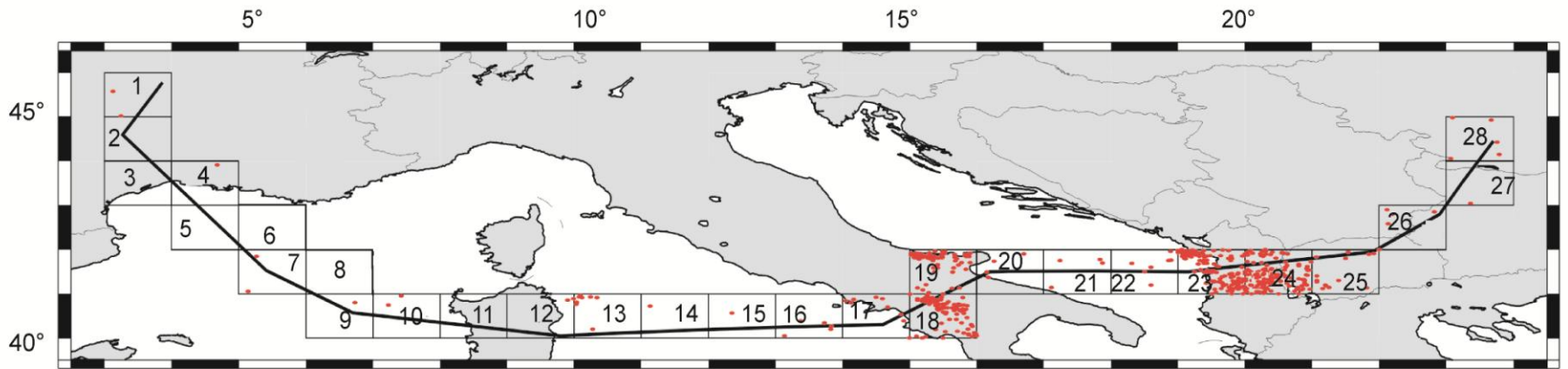
Figure 1

Geologic CO₂ degassing in Italy

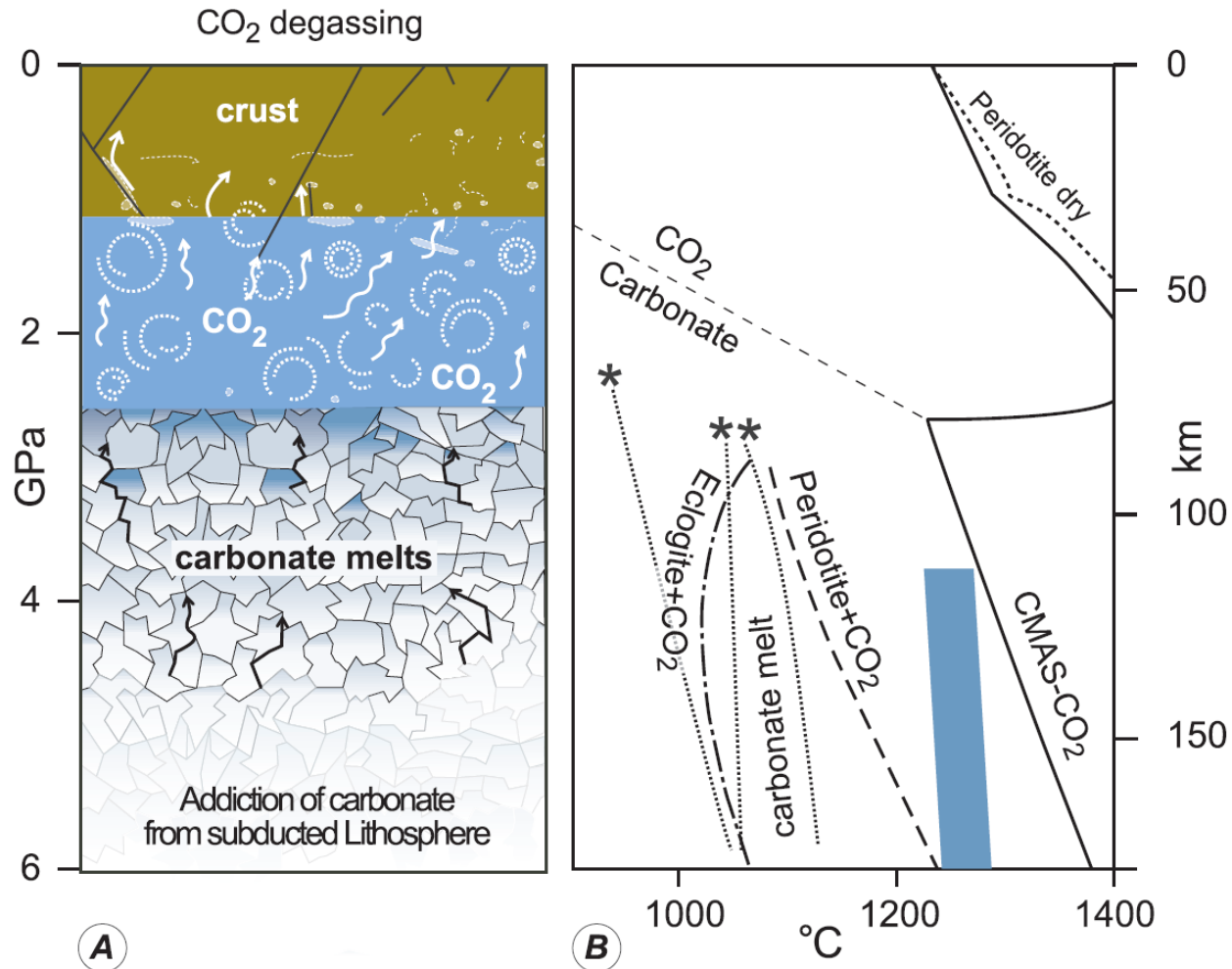
Volcanic		Non-volcanic	
Output (Mt/y)		Output (Mt/y)	
Crater emission		Regional	
Etna (1976 - 1985) ^{1, 2}	25.5	Central Italy ¹³	> 4
Etna (1993 - 1997) ³	4 - 13	Central Apennines ¹⁴	4 - 13.2
Stromboli ⁴	1 - 2	Tuscany and N Latium	6
Vulcano ⁵	0.066	Campania ¹⁵	3
Ground emission		Central Italy ¹⁶	
Vulcano Fossa crater ⁶	0.073	9.7 - 17	
Vulcano plains ⁷	0.027	Soil degassing	
Vulcano fumaroles ⁸	0.088	Latera, Vulsini ¹³	> 0.07
Stromboli ⁹	0.07 - 0.09	Alban Hills ¹⁶	0.2
Vesuvio ¹⁰	0.5	Siena graben ¹⁷	> 0.5
Solfatara, Phl. Fields ¹¹	1.8	Ustica ¹⁸	0.02
Ischia ¹²	0.14	Gas vents	
		Mefite d'Ansanto ¹³	0.3
		Rapolano, Tuscany ¹⁵	0.035
		Mofeta dei Palici, Sicily ¹⁹	0.091
		Geothermal Fields	
		Mt. Amiata ²⁰	0.5

- 1) Allard et al., 1991; 2) Gerlach, 1991; 3) Allard, et al., 1998; 4) Allard et al., 1994; 5) Baubron et al., 1990
 6) Chiodini et al., 1996; 7) Chiodini et al., 1997; 8) Italiano et al., 1998; 9) Carapezza and Federico, 2000
 10) Frondini et al., 2004; 11) Caliro et al., 2008; 12) Aiuppa et al., 2007; 13) Rogie et al., 2000;
 14) Chiodini et al., 2000; 15) Chiodini et al., 2004; 16) Gambardella et al., 2004; 17) Etiope, 1995;
 18) Etiope et al., 1999; 19) Di Gregorio et al., 2002; 20) Frondini et al., 2008

Section 1



Carbon degassing beneath western Mediterranean region



Deep Carbon cycling beneath western Mediterranean region

Volcanic degassing:

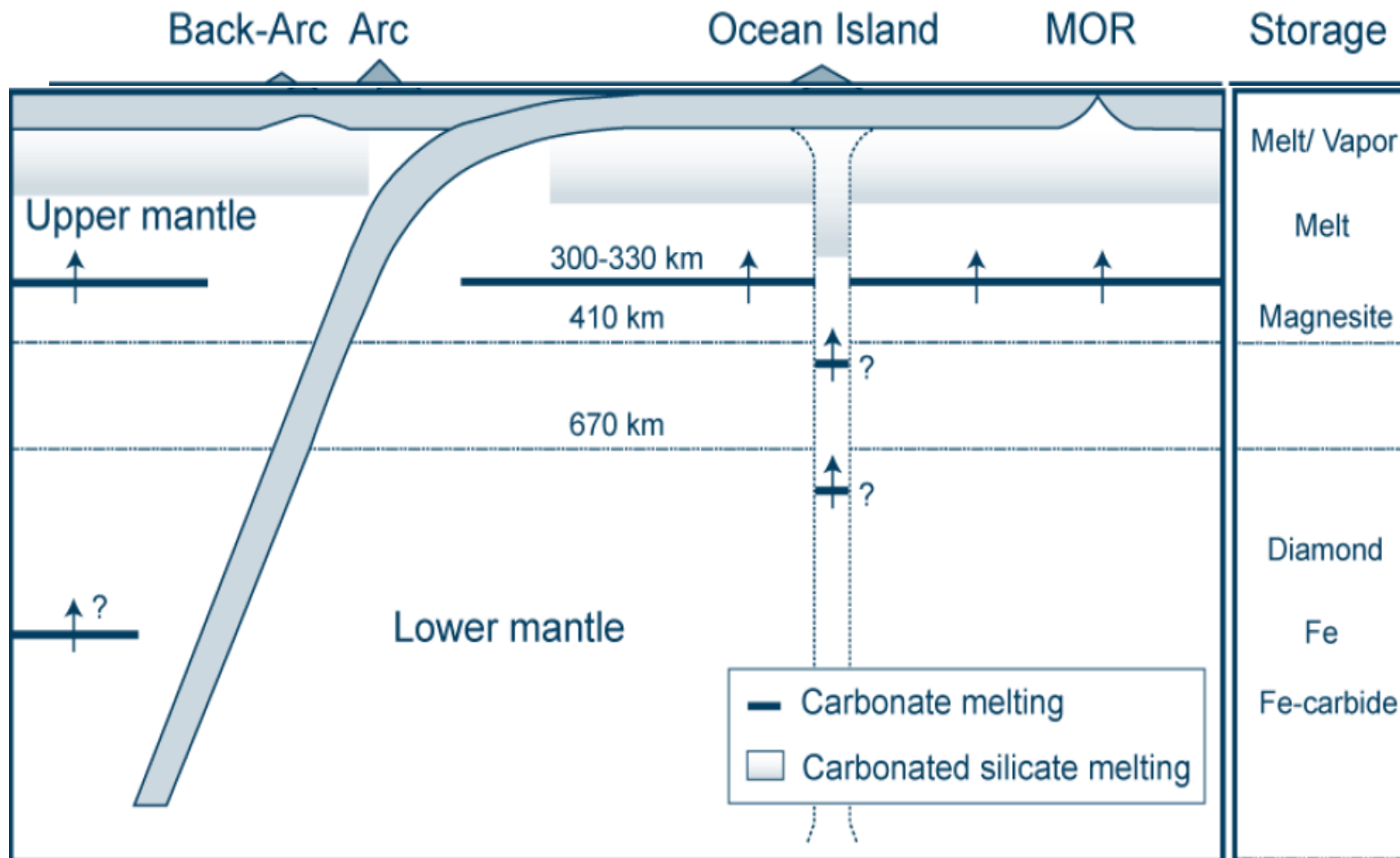
- 20 -30 Mt/yr

Non-volcanic degassing:

- 15 – 30 Mt/yr

- Carbonate-rich melts 45% CO₂ by weight.
- 0.1 - 1 wt% of carbonate melt concentration, and 100% degassing=1.35 – 13.5 Mt of CO₂ (equal to 0.4 - 4 Mt carbon) for each km³ of metasomatized mantle.
- Assuming 30 M.y., CO₂ degassing of the low velocity wedge beneath the Western Mediterranean would conservatively lead to lithosphere-asthenosphere CO₂ flux of about 60 - 600 Mt/yr

Deep Carbon cycling



Conclusions

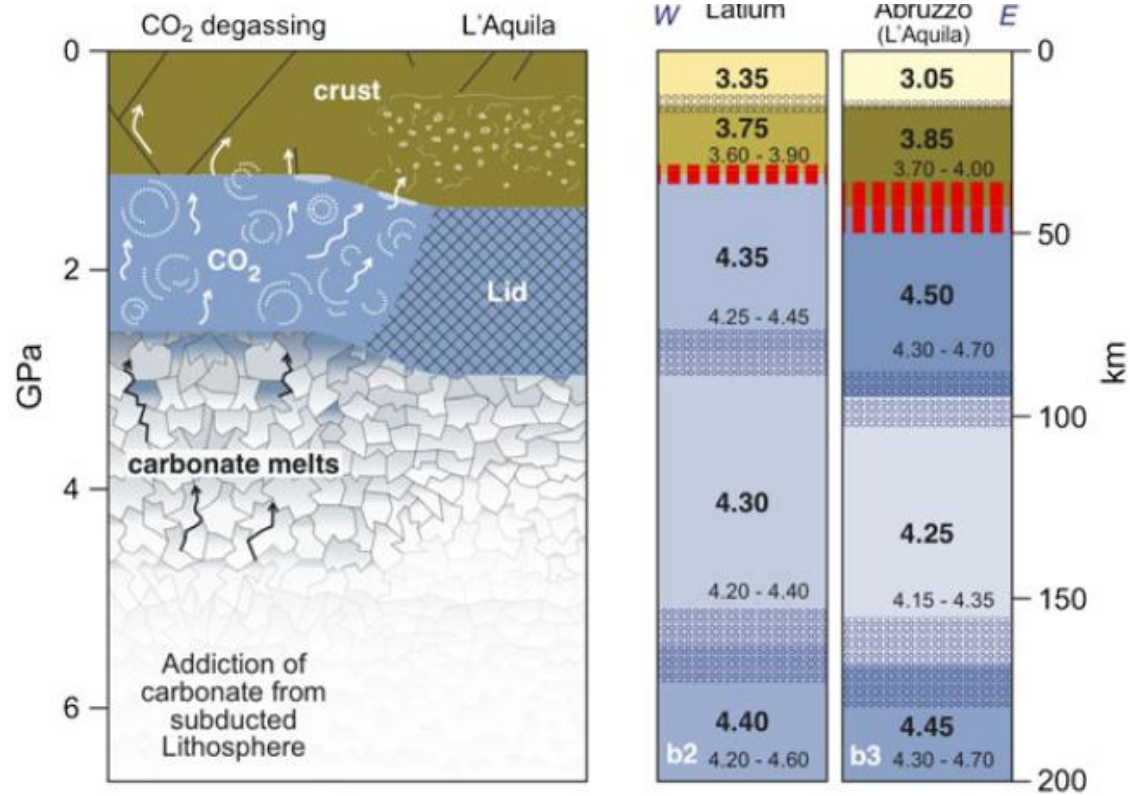


The long-term fluxes in the carbon cycle control the Earth's long-term climate.

Petrological modelling of carbon cycling suggests that carbonated rock melting regulate the deep carbon fluxes from the mantle at the global scale

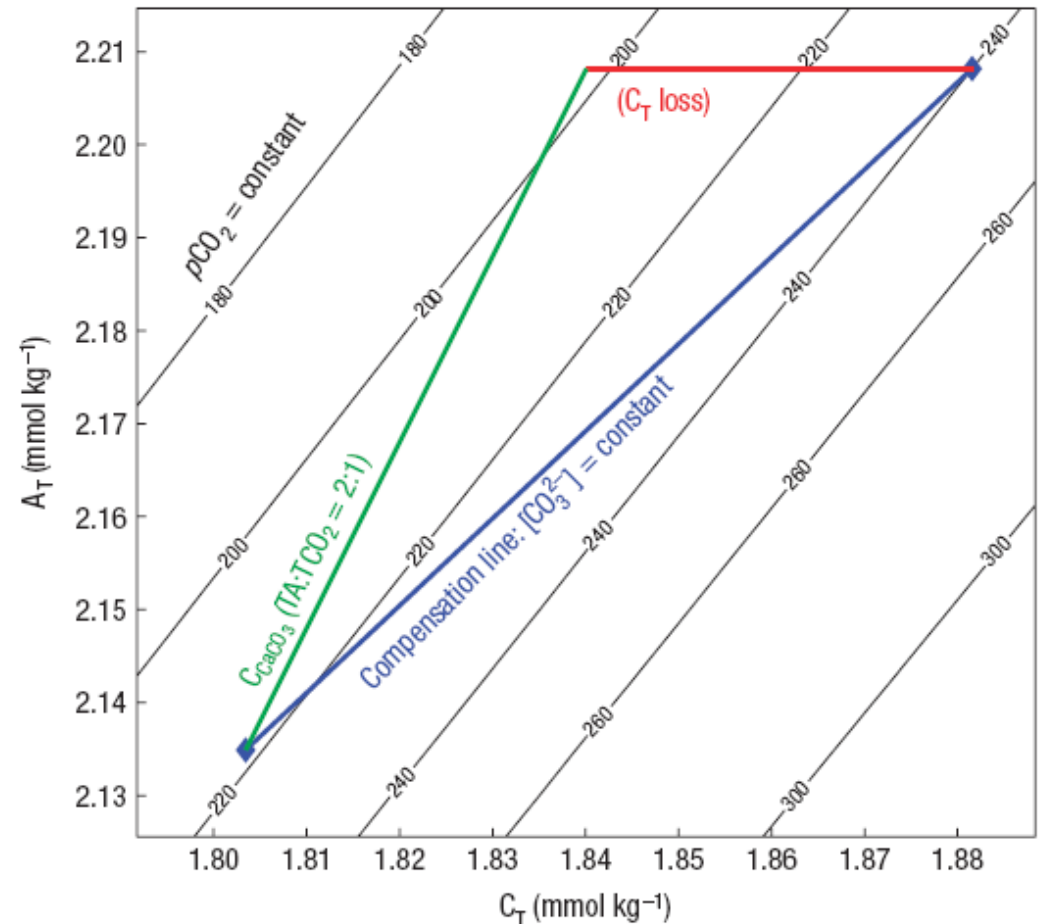
The introduction of other variables (e.g. fO_2) will improve modeling deep carbon cycling





Imbalance of long-term Carbon fluxes

- Imbalance 1-2% past 610 kyr
- -22 to +10 ppm CO₂ in atmosphere
- Carbonate compensation (10 kyr)





- Camerino, 18 Gennaio (Dipartimento di Scienze della Terra);
Urbino, 19 gennaio (Scuola di Dottorato in Scienze della Terra);
Siena, 26 gennaio (Scuola di Dottorato in Scienze della Terra-Preistoria);
Padova, 2 febbraio (Dipartimento di Geoscienze);
Parma, 15 febbraio (Dipartimento di Scienze della Terra);
Genova, 16 febbraio (Scuola di Dottorato STAT);
Torino, 23 febbraio (Dipartimento di Scienze Mineralogiche e Petrologiche);
San Marco in Lamis (FG), 1 marzo (Museo del Carsismo e dei Dinosauri).